



Emissions of methane in Europe inferred by total column measurements

Debra Wunch¹, Dylan B. A. Jones¹, Geoffrey C. Toon², Nicholas M. Deutscher^{3,4}, Frank Hase⁵, Justus Notholt⁴, Ralf Sussmann⁶, Thorsten Warneke⁴, Jeroen Kuenen⁷, Hugo Denier van der Gon⁷, Jenny A. Fisher³, and Joannes D. Maasakkers⁸

¹Department of Physics, University of Toronto, Toronto, Ontario, Canada

²Jet Propulsion Laboratory, California Institute of Technology, Pasadena, California, USA

³Centre for Atmospheric Chemistry, University of Wollongong, Wollongong, New South Wales, Australia

⁴Institute of Environmental Physics, University of Bremen, Germany

⁵Karlsruhe Institute of Technology, IMK-ASF, Karlsruhe, Germany

⁶Karlsruhe Institute of Technology, IMK-IFU, Garmisch-Partenkirchen, Germany

⁷TNO Dept Climate, Air and Sustainability, Utrecht, the Netherlands

⁸School of Engineering and Applied Sciences, Harvard University, Cambridge, Massachusetts, USA

Correspondence: Debra Wunch (dwunch@atmosph.physics.utoronto.ca)

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Abstract. Using five long-running ground-based atmospheric observatories in Europe, we demonstrate the utility of long-term, stationary, ground-based measurements of atmospheric total columns for verifying annual methane emission inventories. Our results indicate that the methane emissions for the region in Europe between Orléans, Bremen, Białystok, and Garmisch-Partenkirchen are overestimated by the state-of-the-art inventories of the Emissions Database for Global Atmospheric Research (EDGAR) v4.2 FT2010 and the high-resolution emissions database developed by the Netherlands Organisation for Applied Scientific Research (TNO) as part of the Monitoring Atmospheric Composition and Climate project (TNO-MACC_III), possibly due to the disaggregation of emissions onto a spatial grid. Uncertainties in the carbon monoxide inventories used to compute the methane emissions contribute to the discrepancy between our inferred emissions and those from the inventories.

1 Introduction

Recent global policy agreements have led to renewed efforts to reduce greenhouse gas emissions to cap global temperature rise (e.g., Conference of the Parties 21, COP 21; UN-

FCCC, 2015; Kona et al., 2016). This, in turn, has motivated countries to seek methods of reducing their greenhouse gas emissions. In Europe, methane emissions account for a significant fraction (about 11 % by mass of CO₂ equivalent) of the total greenhouse gas emissions (UNFCCC, 2017). The lifetime of atmospheric methane is significantly shorter than for carbon dioxide, its 100-year global warming potential is significantly larger, and it is at near steady state in the atmosphere; therefore, significant reductions in methane emissions are an effective short-term strategy for reducing greenhouse gas emissions (Dlugokencky et al., 2011). Emission reduction strategies that include both methane emission reductions and carbon dioxide reductions are thought to be among the most effective at slowing the increase in global temperatures (Shoemaker et al., 2013). Thus, it is important to know exactly how much methane is being emitted and the geographic and temporal source of the emissions. This requires an approach that combines state-of-the-art emissions inventories that contain information about the specific point and area sources of the known emissions and timely and long-term measurements of greenhouse gases in the atmosphere to verify that the emissions reduction targets are met.

Because atmospheric methane is well-mixed and has a lifetime of about 12 years (Stocker et al., 2013), it is trans-

ported far from its emission source, making source attribution efforts challenging from atmospheric measurements alone. Atmospheric measurements are often assimilated into “flux inversion” models to locate the sources of the emissions (e.g., Houweling et al., 2014) but rely on model wind fields to drive transport, and they also tend to have spatial resolutions that do not resolve subregional scales. Methane measurement schemes that constrain emissions on local and regional scales are thus important to help identify the sources of the emissions and to verify inventory analyses. Regional- or national-scale emissions are important to public policy as those emissions are reported annually to the United Nations Framework Convention on Climate Change (UNFCCC).

The atmospheric measurement techniques that are used to estimate methane emissions include measurements made in situ, either on the ground, from tall towers, or from aircraft. Remote sensing techniques are also used, either from space or from the ground. The spatial scale of the sensitivity to emissions differs with the measurement technique: surface in situ measurements provide information about local emissions on urban scales (e.g., McKain et al., 2015; Hopkins et al., 2016), and aircraft in situ measurements can provide information about regional- and synoptic-scale fluxes (e.g., Jacob et al., 2003; Kort et al., 2008, 2010; Wofsy, 2011; Baker et al., 2012; Frankenberg et al., 2016; Karion et al., 2016). Satellite remote sensing techniques provide information useful for extracting emission information on larger scales (regional to global) (e.g., Silva et al., 2013; Schneising et al., 2014; Alexe et al., 2015; Turner et al., 2015) and for large point or urban sources (e.g., Kort et al., 2012, 2014; Nassar et al., 2017). Several studies have shown the importance of simultaneous measurements of co-emitted species (e.g., C_2H_6 and CH_4 or CO and CO_2 , Aydin et al., 2011; Simpson et al., 2012; Peischl et al., 2013; Silva et al., 2013; Hausmann et al., 2016; Wunch et al., 2016; Jeong et al., 2017) or co-located measurements (e.g., Wunch et al., 2009, 2016), showing the added analytical power of the combination of atmospheric tracer information. Ground-based remote sensing instruments have been used to estimate methane emissions on urban (e.g., Wunch et al., 2009; Hase et al., 2015; Wunch et al., 2016) and sub-urban (e.g., Chen et al., 2016; Viatte et al., 2017) scales. In Hase et al. (2015), Viatte et al. (2017), and Chen et al. (2016), the authors have placed mobile ground-based remote sensing instruments around a particular emitter of interest (e.g., a city, dairy, or neighborhood) and have designed short-term campaigns to measure the difference between upwind and downwind atmospheric methane abundances. From these differences the authors have computed emission fluxes. However, there is a network of nonmobile ground-based remote sensing instruments that have been collecting long-term measurements of atmospheric greenhouse gas abundances. These instruments were not placed intentionally around an emitter of interest, but collectively they ought to contain information about nearby emissions. To date, there have been no studies that have at-

tempted to extract regional methane emission information from these existing ground-based remote sensing observatories.

In this paper, we will describe our methods for computing the emissions of methane using five stationary ground-based remote sensing instruments located in Europe in Sect. 2. Our results and comparisons to the state-of-the-art inventories are shown in Sect. 3, and we summarize our results in Sect. 4.

2 Methods

Our study area is the region between five long-running atmospheric observatories situated in Europe. Three of the stations are in Germany: Bremen (Notholt et al., 2014), Karlsruhe (Hase et al., 2014), and Garmisch-Partenkirchen (Sussmann and Rettinger, 2014). The other two are in Poland (Białystok, Deutscher et al., 2017) and France (Orléans, Warneke et al., 2014). Each station measures the vertical column-averaged dry-air mole fraction of carbon dioxide (X_{CO_2}), carbon monoxide (X_{CO}), methane (X_{CH_4}), and other trace gas species. The locations are shown in Fig. 1, overlaid on a nighttime light image from the National Aeronautics and Space Administration (NASA) to provide a sense of the population density of the area. These observatories are part of the Total Carbon Column Observing Network (TCCON, Wunch et al., 2011) and have been tied to the World Meteorological Organization trace-gas scale through comparisons with vertically integrated, calibrated in situ profiles over the observatories (Wunch et al., 2010; Wunch et al., 2015; Messerschmidt et al., 2011; Geibel et al., 2012).

Following a similar method to Wunch et al. (2009, 2016), we estimate emissions of methane from the data recorded from the TCCON observatories, coupled with gridded inventories of carbon monoxide within the region. We compute changes (or “anomalies”) in X_{CH_4} and X_{CO} that we will refer to as ΔX_{CH_4} and ΔX_{CO} , and we then compute the slopes relating ΔX_{CH_4} to ΔX_{CO} . From the computed slopes (α), we can infer emissions of methane (E_{CH_4}) if emissions of carbon monoxide (E_{CO} , in mass per unit time) are known, using the following relationship:

$$E_{CH_4} = \alpha \frac{m_{CH_4}}{m_{CO}} E_{CO}, \quad (1)$$

where $\frac{m_{CH_4}}{m_{CO}}$ is the ratio of the molecular masses of CH_4 and CO .

In Wunch et al. (2009, 2016), measurements from a single atmospheric observatory were used to infer emissions because the unique dynamics of the region advected the polluted air mass into and out of the study area diurnally. In this paper, we rely on several stations to provide measurements of the boundary of the study region to measure CO and CH_4 emitted between the stations. This analysis relies on a few assumptions about the nature of the emissions. First, that the lifetimes of the gases of interest are longer than the transport



Figure 1. The locations of the TCCON observatories overlaid on a NASA nighttime light image. From west to east, the stations are Orléans (or, pink), Karlsruhe (ka, green), Bremen (br, blue-green), Garmisch-Partenkirchen (gm, orange), and Białystok (bi, purple).

time within the region. This is the case both for methane, which has an atmospheric lifetime of 12 years, and for carbon monoxide, which has an atmospheric lifetime of a few weeks. Second, we assume that typical emissions are consistent over time periods longer than a few days so that they are advected together. The nature of the emissions in this region (mostly residential and industrial energy needs) supports this assumption. Third, we assume that the spatial distribution of the emissions is similar for CH_4 and CO , as confirmed by the inventory maps (Fig. A3). This method does not require carbon monoxide and methane to be co-emitted (as they generally do not have the same emissions sources).

To compute anomalies and slopes, we first filter the data to minimize the impact of data sparsity and air mass differences between stations (Appendix A). Then, for each station, the daily median value is subtracted from each measurement. This reduces the impact of the station altitude and any background seasonal cycle from aliasing into the results. Subsequently, we compute the differences in the X_{CH_4} and X_{CO} abundances measured at the same solar zenith and solar azimuth angles on the same day at two TCCON stations. By computing anomalies at the same solar zenith angles, we minimize any impact that air-mass-dependent biases could have on the calculated anomalies. This analysis is repeated for all combinations of pairs of stations within the study area. The vertical sensitivity of the TCCON measurements is explicitly taken into account by dividing the anomalies by the

surface layer column averaging kernel value, as we assume that the anomalies are due to emissions near the surface. The slopes computed for each year and each pair of stations are shown in Fig. 2.

The farthest distance between the European TCCON stations included in this study is between Orléans and Białystok (1580 km). Climatological annual mean surface wind speeds from the National Centers for Environmental Prediction (NCEP) and National Center for Atmospheric Research (NCAR) reanalysis (Kalnay et al., 1996) within the study area are about 6 km h^{-1} (Fig. A1). The air from Orléans will quickly mix vertically from the surface where the winds aloft are more rapid than at the surface (see Appendix B). Thus, air from Orléans would normally reach Białystok in a few days. To determine whether these anomalies are consistent throughout the transport time through the study area, we compute anomalies between sites lagged by up to 14 days. The slopes of the anomalies do not change significantly or systematically with the lag time (Appendix B; Fig. A2), presumably because the atmospheric composition within the study area is relatively well-mixed or because the emissions are relatively consistent from day to day within the study area.

Previous papers have used carbon dioxide instead of carbon monoxide to infer methane emissions. We choose to compute emissions using measurements of X_{CO} instead of X_{CO_2} in this work because the natural CO_2 fluxes in the re-

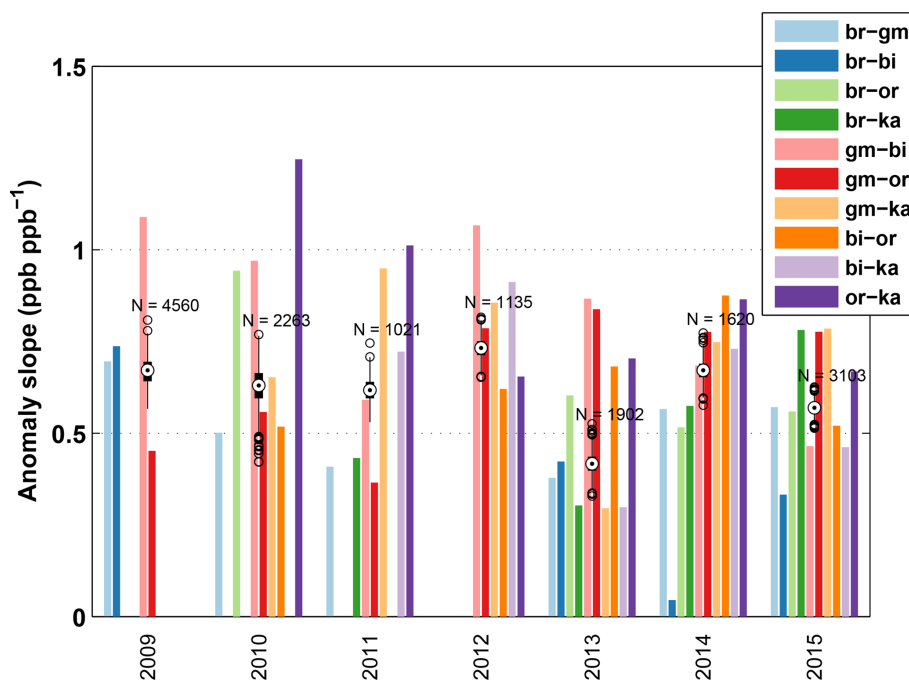


Figure 2. The bars show the methane to carbon monoxide anomaly slopes for each site pair. The method of computing these anomaly slopes is detailed in Sect. 2 of the main text. The black targets indicate the median value of the slope for that year, when all site pairs are considered simultaneously, and the 25th and 75th quartiles of the median value are indicated by the vertical black bars. Outliers are indicated by open black circles.

gion are large compared with the anthropogenic emissions, and they have a strong diurnal and seasonal cycle. The distance between the stations is large enough that local (sub-daily) uptake of CO_2 differs from station to station, significantly obscuring the relationships between methane and carbon dioxide, and thus the anomaly slopes, especially in the summer months. While the emissions inventory of anthropogenic CO_2 may be more accurate than the CO inventory in the region, the presence of these large natural fluxes of CO_2 precludes its use in the anomaly slope calculation. The accuracy of our method, therefore, is limited by the accuracy of the carbon monoxide emission inventory. Fires could provide a large flux of CO without a large CH_4 flux, and this should also be taken into consideration in these types of analyses. In our study area fluxes from fires are small.

2.1 Inventories

To obtain an estimate of carbon monoxide emissions (E_{CO}) within the study area, we use gridded inventories and sum the emissions within the study area to compare with our emissions inferred from the TCCON measurements (see Appendix C and Fig. A3 for details). The two inventories employed here are the Emissions Database for Global Atmospheric Research (EDGAR) and the Netherlands Organisation for Applied Scientific Research (TNO) high-resolution emissions database developed as part of the Monitoring Atmospheric Composition and Climate project

(TNO-MACC_III). The EDGAR version v4.3.1_v2 of January 2016 annual gridded inventory is available at $0.1^\circ \times 0.1^\circ$ spatial resolution and reports global emissions from the year 2000 to 2010 (Olivier et al., 1994; EC-JRC and PBL, 2016). The TNO-MACC_III inventory is a Europe-specific air quality emissions inventory, available on a $0.125^\circ \times 0.0625^\circ$ grid, and reports emissions for 2000–2011 (Kuenen et al., 2014). Both EDGAR and TNO-MACC_III provide spatially and temporally coincident methane inventories which we use to compare with our inferred emissions. We use the EDGAR version v4.2 FT2010 and the TNO-MACC_III methane inventories.

Using country-level emissions reported through 2015 from the European Environment Agency (EEA, 2015), we extrapolate the EDGAR and TNO-MACC_III gridded inventory CO emissions for the study area through 2015. This facilitates more direct comparisons with the TCCON measurements, which begin with sufficient data for our study in 2009. We extrapolate the emissions by scaling the total emissions from the countries that are intersected by the area of interest (Germany, Poland, Belgium, France, Luxembourg, and the Czech Republic) to the last reported year of emissions from the inventory. We then assume that the same scaling factor applies for each subsequent year. The details of the extrapolation method are in Appendix D and Figs. A4 and A5.

The time series of the reported emissions from 2000 to 2015 are shown in Fig. 3. The inventories and scaled

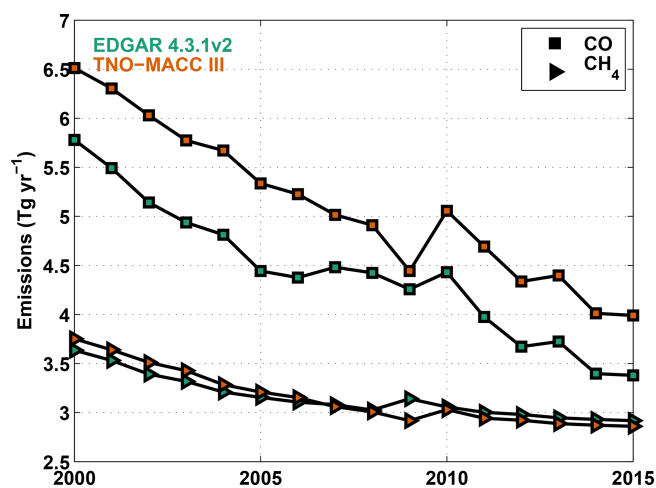


Figure 3. This figure shows the summed EDGAR (green) and TNO-MACC_III (orange) emissions within the study area for CO (squares) and CH₄ (triangles). The study area is defined in Fig. 1. All emissions are shown in units of Tg yr⁻¹. Extrapolation begins after 2010 for EDGAR and 2011 for TNO-MACC_III.

country-level reported emissions for this region suggest that emissions of CO and CH₄ have decreased by about 40 % and 20 %, respectively, between 2000 and 2015. The TNO-MACC_III carbon monoxide emissions are on average 15 % higher than the EDGAR v4.3.1 emissions in the study area. The total TNO-MACC_III and EDGAR methane emissions agree to within 2 % in the study area.

An earlier version of the EDGAR carbon monoxide inventory was evaluated by Stavrou and Müller (2006) and Fortems-Cheiney et al. (2009), who assimilated satellite measurements of CO using the EDGAR v3.3FT2000 CO emissions inventory as the a priori. Stavrou and Müller (2006) found that, over Europe, the a posteriori emissions increase by less than 15 % when assimilating carbon monoxide from the Measurements of Pollution in the Troposphere (MOPITT) satellite instrument (Emmons et al., 2004). Fortems-Cheiney et al. (2009) assimilated Infrared Atmospheric Sounding Interferometer (IASI) CO (Clerbaux et al., 2009) and MOPITT CO and found that the a posteriori emissions increase by 16 % and 45 %, respectively.

The more recent EDGAR v4.3.1 CO emissions in our study are 24 % lower than the EDGAR v3.3FT2000 CO emissions for the year 2000, so it may be that the EDGAR v4.3.1 CO emissions are significantly underestimated. However, assimilations of CO are known to be very sensitive to the chemistry described in the model: most notably the OH chemistry (Protonotariou et al., 2010; Yin et al., 2015). Therefore, it is difficult to determine how much of the discrepancy between versions of the model is from the inventory or the model chemistry.

The EDGAR methane inventory has been evaluated in several previous studies. It has been shown to overestimate re-

gional CH₄ emissions (e.g., Wunch et al., 2009; Wecht et al., 2014) but to underestimate oil and gas emissions (e.g., Miller et al., 2013; Buchwitz et al., 2017). However, recent methane isotope analysis by Röckmann et al. (2016) has suggested that the EDGAR inventory overestimates fossil-fuel-related emissions. The study area of interest here has little oil and gas production, except for some test sites in Poland (USEIA, 2015), no commercial shale gas industry, and few pipelines.

2.2 Model experiment

To test whether the anomaly method described in Sect. 2 can accurately infer methane emissions, we conducted a modeling experiment using version v12.1.0 of the GEOS-Chem model (<http://www.geos-chem.org>, last access: 4 January 2019) to simulate methane and carbon monoxide for the year 2010. The model is driven by the Modern-Era Retrospective analysis for Research and Applications, version 2 (MERRA-2) meteorology from the NASA Global Modeling and Assimilation Office. The native resolution of the meteorological fields is 0.25° × 0.3125°, with 72 vertical levels from the surface to 0.01 hPa, which we degraded to 2° × 2.5° and 47 vertical levels. We use the linear CO-only and CH₄-only simulations of GEOS-Chem, with prescribed monthly mean OH fields. In the CO-only simulation, global anthropogenic emissions are from EDGAR v4.3.1, which are overwritten regionally with the following emissions: the Cooperative Programme for Monitoring and Evaluation of the Long-range Transmission of Air Pollutants in Europe (EMEP), the U.S. Environmental Protection Agency National Emission Inventory for 2011 (NEI2011), the MIX inventory for Asia, the Visibility Observational (BRAVO) Study Emissions Inventory for Mexico, and the criteria air contaminants (CAC) inventory for Canada. The sources of CO from the oxidation of CH₄ and volatile organic compounds (VOCs) are prescribed following Fisher et al. (2017). For the CH₄-only simulation, the emissions are as described in Maasakkers et al. (2019). Global anthropogenic emissions are from EDGAR v4.3.2, but the US emissions were replaced with those from Maasakkers et al. (2016), and emissions from wetlands are from WetCHARTs version 1.0 (Bloom et al., 2017). For both CO and CH₄ simulations, emissions from biomass burning are from the Quick Fire Emissions Dataset (QFED) (Darmenov and Silva, 2015). The biomass burning in the study area produces less than 2 % of the total anthropogenic emissions of CO.

We used identical OH fields (from version v7-02-03 of GEOS-Chem) for the CO and CH₄ simulations, so that the chemical losses of methane and carbon monoxide are consistent, and ran tagged CO experiments so that we could identify the source of the emissions. The model atmospheric carbon monoxide and methane profiles were integrated to compute simulated X_{CO} and X_{CH₄}. To illustrate the sensitivity of the modeled fields to European emissions, we show the seasonal means of the modeled X_{CO} sampled at the five TC-

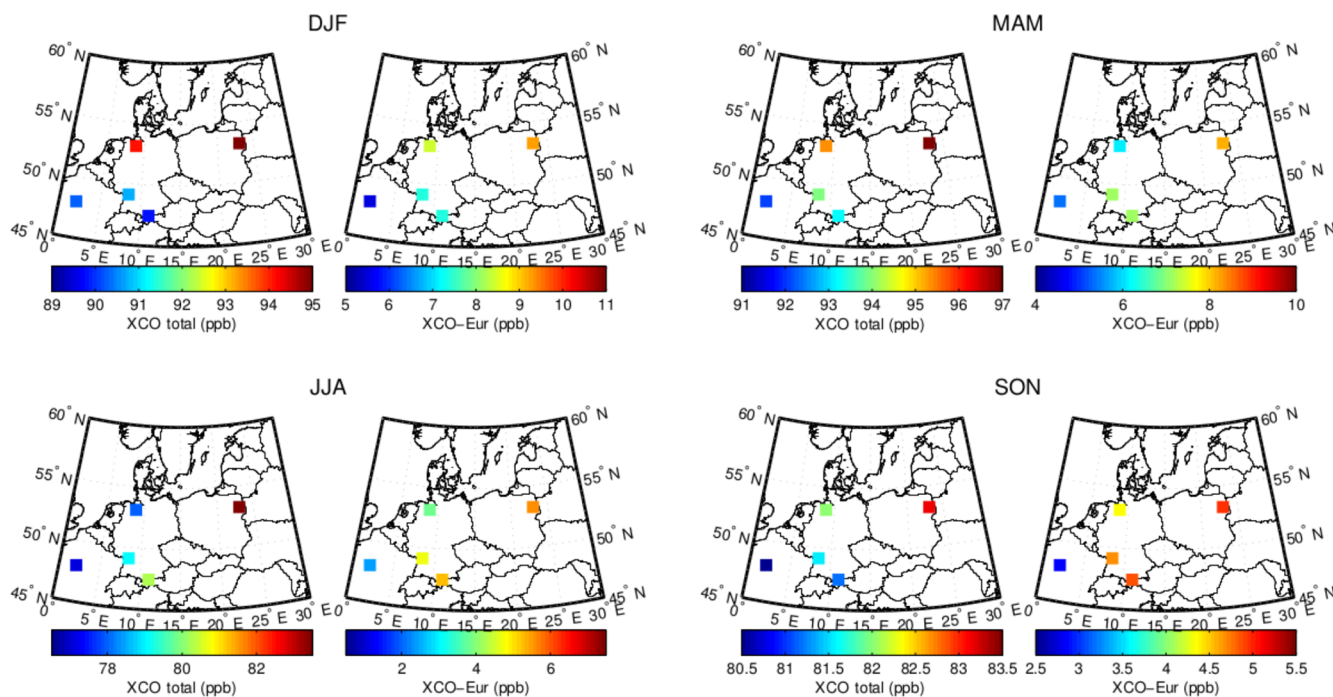


Figure 4. This figure compares seasonally averaged modeled total X_{CO} with the X_{CO} contribution from emissions in Europe. Each season has two maps: the left map shows the total X_{CO} and the right map shows the contribution from European emissions (X_{CO-Eur}). The spatial pattern of the gradients in modeled X_{CO} between the TCCON stations is reflected in the European contribution.

CON stations in Fig. 4. Also plotted is the column contribution (X_{CO-Eur}) from CO emissions only in Europe (defined as the broader region between $0-45^{\circ}$ E and $45-55^{\circ}$ N). As can be seen, the spatial pattern of the differences in modeled X_{CO} between the TCCON stations is reflected in X_{CO-Eur} . We calculated the anomalies in X_{CO} and X_{CO-Eur} , using the same approach employed with the atmospheric data, and found that the anomalies in X_{CO-Eur} , which represent the direct influence of European emissions on atmospheric CO, account for about 35% of the anomalies in X_{CO} . This confirms that the X_{CO} anomalies between the TCCON stations are sensitive to European emissions.

To estimate the modeled CH_4 emissions using the modeled CO, the modeled X_{CO} and X_{CH_4} were interpolated to the locations of the TCCON stations and anomalies and slopes were computed. We then applied Eq. (1) to our anomaly slopes to compute methane emissions from the known CO emissions, accounting for only the CO emissions from anthropogenic, biomass burning, and biofuel sources. We neglect sources of CO emissions from the oxidation of CH_4 and VOCs because the column enhancements for those emissions are relatively spatially uniform across this region of Europe, and thus they should not contribute significantly to the anomalies. The resulting annual CH_4 emissions agree well with the model emissions: the inferred emissions from the anomaly analysis are higher than the model emissions by less than 2% (Fig. 5).

While the inferred annual emissions agree well with the modeled annual emissions, the seasonal pattern of the emissions inferred from the anomaly analysis differs from that of the model. The anomaly analysis overestimates emissions in the winter and underestimates emissions in the summer. This may be due to small spatial inhomogeneities in the column enhancements from VOC (biogenic) emissions that influence the anomaly analysis most in summertime when VOC emissions are largest. Including the VOC emissions in the total carbon monoxide emissions leads us to infer annual methane emissions that are overestimated by 15%, increasing the inferred summertime emissions without significantly changing the inferred wintertime emissions.

The seasonal analysis suggests that the 2% agreement in the annual emission estimate may reflect the compensating effects of discrepancies over the seasonal cycle, and improving the seasonal estimate may require a better treatment of the VOC contribution to atmospheric CO. Nevertheless, the results here suggest that for this region of Europe, where VOC and methane oxidation emissions lead to relatively spatially uniform column enhancements and fire emissions are small, we can successfully use the anomaly method described in Sect. 2 to infer annual methane emissions.

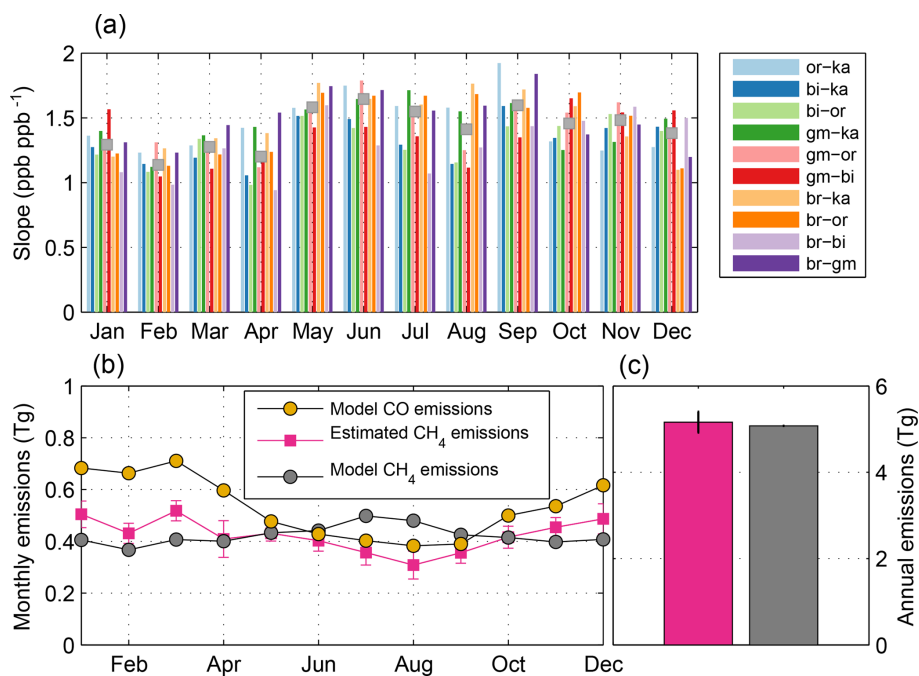


Figure 5. This figure shows the results from the modeling experiment using GEOS-Chem. Panel (a) shows the model $\Delta X_{\text{CH}_4} - \Delta X_{\text{CO}}$ slopes for each month and pair of stations (indicated by the colors). The median slopes for each month are overlaid with grey squares. Panel (b) shows the model carbon monoxide emissions (excluding VOC and methane oxidation) and the model methane emissions. The inferred methane emissions from our tracer–tracer slope method are plotted in pink squares. Panel (c) shows the annual methane emissions from the tracer–tracer slope method and the model.

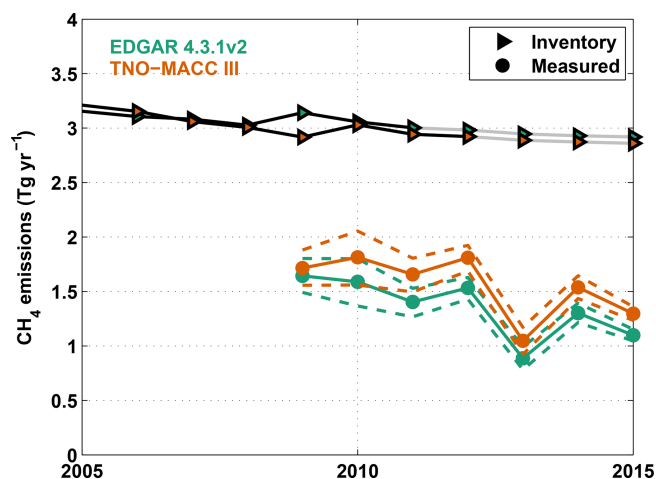


Figure 6. The black line is the summed EDGAR (green) and TNO-MACC_III (orange) methane emissions within the study area shown in Fig. 1. The grey lines indicate the projected emissions based on scaling the country-level emissions reported by the UNFCCC (UNFCCC, 2017) to the area emissions in 2010 for EDGAR and 2011 for TNO-MACC_III. The lower solid lines show the emissions inferred from the TCCON anomaly analysis using CO emissions from the two models, and the dashed lines indicate the 5th and 95th percentiles.

3 Results and discussion

To compute methane emissions, we apply Eq. (1) to our anomaly slopes and the inventory-reported carbon monoxide emissions in the study region (Fig. 6). If we choose the mean of the reported CO emissions from EDGAR v4.3.1 and TNO-MACC_III, the methane emissions we compute within the study area based on the TCCON measurements are $1.7 \pm 0.3 \text{ Tg yr}^{-1}$ in 2009, with a non-monotonic decrease to $1.2 \pm 0.3 \text{ Tg yr}^{-1}$ in 2015. The uncertainties quoted here are from the standard errors on the data slope fitting only; we have not included uncertainties from the inventories. The magnitude of methane emissions we compute from the TCCON data are, on average, about 2.3 times lower than the methane emissions reported by EDGAR and about 2 times lower than the methane emissions reported by TNO-MACC_III.

Our method of inferring methane emissions depends critically on the carbon monoxide inventory. The carbon monoxide emissions for 2010 in the study area from our GEOS-Chem model run, derived from EMEP emissions, were 6.4 Tg , about 35 % higher than the average of the EDGAR and TNO-MACC_III emissions for that year. This magnitude underestimate has also been suggested by Stavrou and Müller (2006) and Fortems-Cheiney et al. (2009) using independent data. Using the GEOS-Chem carbon monoxide

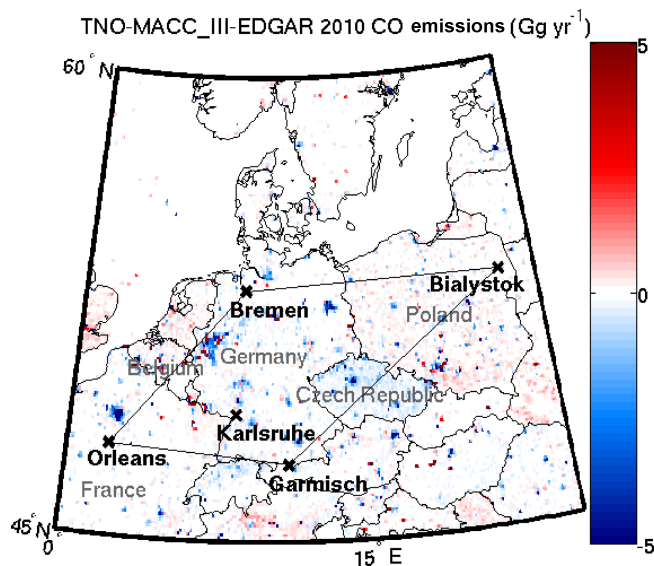


Figure 7. This map shows the difference between the TNO-MACC_III carbon monoxide emissions and the EDGAR emissions for the year 2010. The black straight lines delineate the study area from the surrounding region. The TCCON stations included in this study are marked with black “x” symbols and labeled in black bold font. The countries intersected by or contained within the study area are labeled in grey. Warm (red) colors indicate that the TNO-MACC_III inventory is larger than the EDGAR inventory; cool (blue) colors indicate that the EDGAR inventory is larger than TNO-MACC_III.

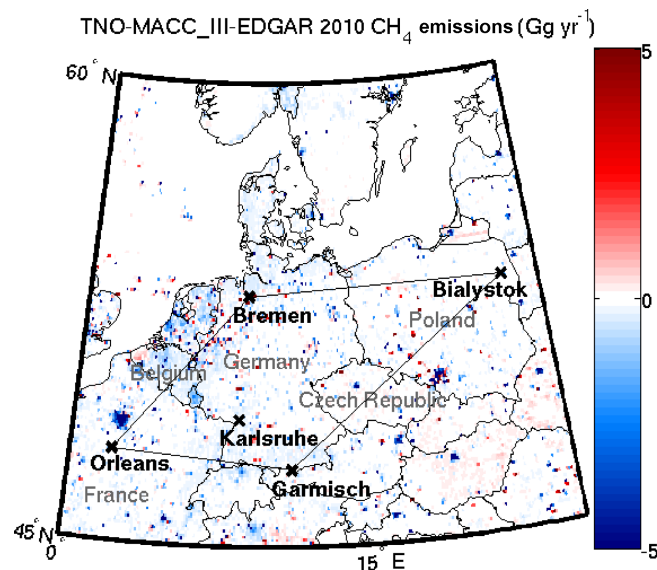


Figure 8. This map shows the difference between the TNO-MACC_III methane emissions and the EDGAR emissions for the year 2010. The labeling and coloring follows that in Fig. 7.

emissions increases the methane emissions inferred by the anomaly analysis to 2.4 ± 0.3 Tg in 2010. This value remains lower than the EDGAR and TNO-MACC_III methane emissions estimates for 2010, which are 3 Tg, but by only 20%. Therefore, we find that the inventories likely overestimate methane emissions, but the accuracy of our results relies on the accuracy of the carbon monoxide inventory.

Although the EDGAR and TNO-MACC_III inventories agree to within 15% in carbon monoxide emissions and 2% in methane emissions in the study region, they spatially distribute these emissions differently. Maps of the spatial differences between the TNO-MACC_III and EDGAR emissions are shown in Fig. 7 for carbon monoxide and Fig. 8 for methane. EDGAR estimates larger emissions of carbon monoxide from the main cities in the study region and the surrounding areas. This is clearly visible from the difference map (Fig. 7), where cities such as Hamburg, Berlin, Prague, Wrocław, Warsaw, Munich, Paris, and Vienna appear in blue. However, the overall carbon monoxide emissions from TNO-MACC_III in the study area are higher than EDGAR, and this comes from regions between the main cities, particularly in Poland and eastern France.

The differences between EDGAR and TNO-MACC_III methane emissions also show that the EDGAR emissions estimates near large cities are significantly larger (Fig. 8).

In contrast to the carbon monoxide spatial distribution, the TNO-MACC_III methane emissions are generally smaller everywhere, except for discrete point sources.

Comparing country-level carbon monoxide emissions reported in 2010 with the inventories shows reasonable agreement, which is expected since the inventories use country-level reports as input. The sum of the carbon monoxide emissions within the entire countries of Germany, Poland, France, Luxembourg, Belgium, and the Czech Republic differ between EDGAR and TNO-MACC_III by 18%, with EDGAR estimates lower than those from TNO-MACC_III. Emissions from Germany, most of which are included in the study area, differ by only 6% between EDGAR and TNO-MACC_III, again with EDGAR estimates lower than TNO-MACC_III. The national carbon monoxide emissions reported to the Convention on Long-range Transboundary Air Pollution (LRTAP Convention, https://www.eea.europa.eu/ds_resolveuid/0156b7a0ca47485593e7754c52c24afd, last access: 15 November 2017, EEA, 2015) agree to within a few percent of the TNO-MACC_III country-level emissions (e.g., 5.5% for Germany in 2010).

The differences between 2010 country-level emissions estimates are larger for methane: EDGAR estimates are larger than TNO-MACC_III estimates by 36% when summing all countries intersected by the study area and 8% when considering only German emissions. The TNO-MACC_III country-level emissions estimates agree to within a few percent of the UNFCCC (http://di.unfccc.int/time_series, last access: 15 November 2017) country-level reported methane emissions (e.g., 8% for Germany in 2010).

The differences between the EDGAR and TNO-MACC_III inventories suggest that the spatial distribution of emissions is less certain than the larger-scale emissions, since the total carbon monoxide and methane emissions between the inventories agree to within 15 % and 2 %, respectively, in the study area, but these estimates can disagree by a factor of 2 on city-level scales.

If we assume that the national-scale methane emissions are correctly reported in EDGAR and TNO-MACC_III, our results indicate that the methane emissions in the region are incorrectly spatially distributed in the inventories. It could be that point or urban sources outside the study area but within the countries intersected by the study area emit a larger proportion of the country-level emissions than previously thought.

4 Conclusions

Using co-located measurements of methane and carbon monoxide from five long-running ground-based atmospheric observing stations, we have shown that in the area of Europe between Orléans, Bremen, Białystok, and Garmisch-Partenkirchen, the inventories likely overestimate methane emissions and point to a large uncertainty in the spatial distribution (i.e., the spatial disaggregation) of country-level emissions. However, the magnitude of our inferred methane emissions relies heavily on the EDGAR v4.3.1 and TNO-MACC_III carbon monoxide inventories, and thus there is a need for rigorous validation of the carbon monoxide inventories.

This study demonstrates the potential of clusters of long-term ground-based stations monitoring total columns of atmospheric greenhouse and tracer gases. It also shows the potential of having co-located measurements of multiple pollutants to derive better estimates of emissions. These types of observing systems can help policy makers verify that greenhouse gas emissions are reducing at a rate necessary to meet regulatory obligations. The atmospheric measurements are agnostic to the source (and country of origin) of the methane, measuring only what is emitted into the atmosphere in a given area. Thus, they can help validate and reveal inadequacies in the current inventories, and, in particular, how country-wide emission reports are disaggregated on a grid. To enhance these results, simultaneous measurements of complementary atmospheric trace gases, such as ethane, acetylene, nitrous oxide, nitrogen dioxide, ammonia, and isotopes, would help distinguish between sources of methane. This would provide additional valuable information that would likely improve inventory disaggregation.

Data availability. TCCON data are available from the TCCON archive, hosted by the California Institute of Technology at <https://tccodata.org>. Karlsruhe data were obtained from

<https://doi.org/10.14291/tcccon.ggg2014.karlsruhe01.R1/1182416> (Hase et al., 2014). Bremen data were obtained from <https://doi.org/10.14291/tcccon.ggg2014.bremen01.R0/1149275> (Notholt et al., 2014). Garmisch data were obtained from <https://doi.org/10.14291/tcccon.ggg2014.garmisch01.R0/1149299> (Sussmann and Rettinger, 2014). Orléans data were obtained from <https://doi.org/10.14291/tcccon.ggg2014.orleans01.R0/1149276> (Warneke et al., 2014). Białystok data were obtained from <https://doi.org/10.14291/tcccon.ggg2014.bialystok01.R1/1183984> (Deutscher et al., 2017). The Emissions Database for Global Atmospheric Research (EDGAR) inventory is available from the European Commission Joint Research Centre (JRC) and the Netherlands Environmental Assessment Agency (PBL), <http://edgar.jrc.ec.europa.eu> (last access: 7 April 2017). The GEOS-Chem v12.1.0 model is available from <https://doi.org/10.5281/zenodo.1553349> (The International GEOS-Chem User Community, 2018).

Appendix A: Filtering

The filtering method was designed to remove days of data for which the atmospheric air mass was inconsistent between sites (e.g., a front was passing through or there were significant stratospheric incursions into the troposphere) and for years in which there were too few simultaneous measurements at a pair of TCCON stations to compute robust annually representative anomalies.

To address the consistency of the air mass between sites, we retained days on which the retrievals of hydrogen fluoride (X_{HF}) were between 50 ppt and 100 ppt, and deviated by less than 10 ppt of the median X_{HF} value for all sites on that day. HF is a trace gas that exists only in the stratosphere, and thus it serves as a tracer of tropopause height (Washenfelder et al., 2003; Saad et al., 2014). Since the concentration of CH_4 decreases significantly above the tropopause in the mid-latitudes, its total column dry-air mole fraction (X_{CH_4}) is sensitive to the tropopause height. Filtering out days on which X_{HF} varies significantly between sites also ensures that the anomalies (and thus the slopes) are minimally impacted by stratospheric variability. This filter removed less than 5% of the data.

To ensure that the anomalies are representative of the full year, we require that each year has 400 coincident measurements across at least three seasons.

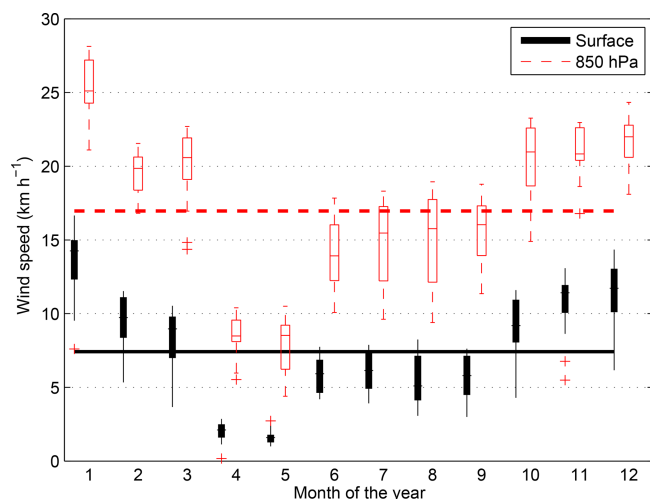


Figure A1. These box plots show the NCEP/NCAR reanalysis long-term climatological monthly mean wind speeds at the surface (filled black boxes) and at 850 hPa (open red boxes) in the study area (see Figs. 1, 7, or 8 for study area maps). The solid black and dashed red horizontal lines indicate the annual mean wind speed at the surface and 850 hPa (~ 1.5 km), respectively. Wind speeds that are aloft (on average 17 km h^{-1}) are significantly swifter than those at the surface (on average 7.5 km h^{-1}).

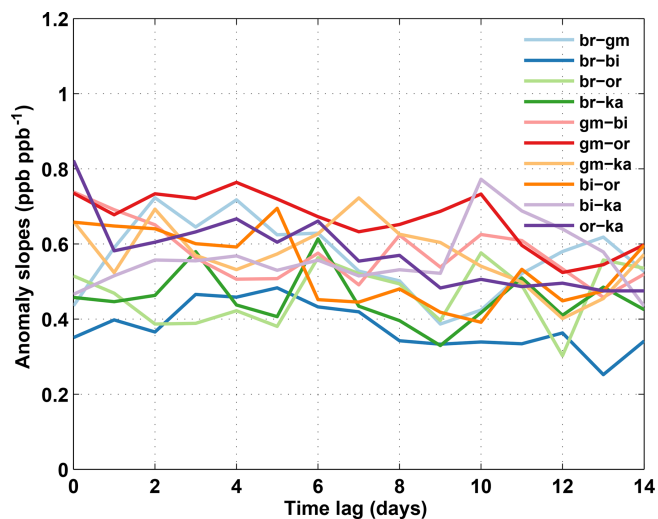


Figure A2. These are the anomaly slopes ($\Delta\text{CH}_4/\Delta\text{CO}$) in ppb ppb^{-1} for each station pair, for the entire time series. The anomalies are computed by subtracting data within the same solar zenith angle bin between two TCCON stations. For more detail, see Sect. 2 of the main text. The x axis indicates the number of days separating the measurements. The legend identifiers are as follows: br – Bremen, gm – Garmisch-Partenkirchen, bi – Białystok, or – Orléans, ka – Karlsruhe.

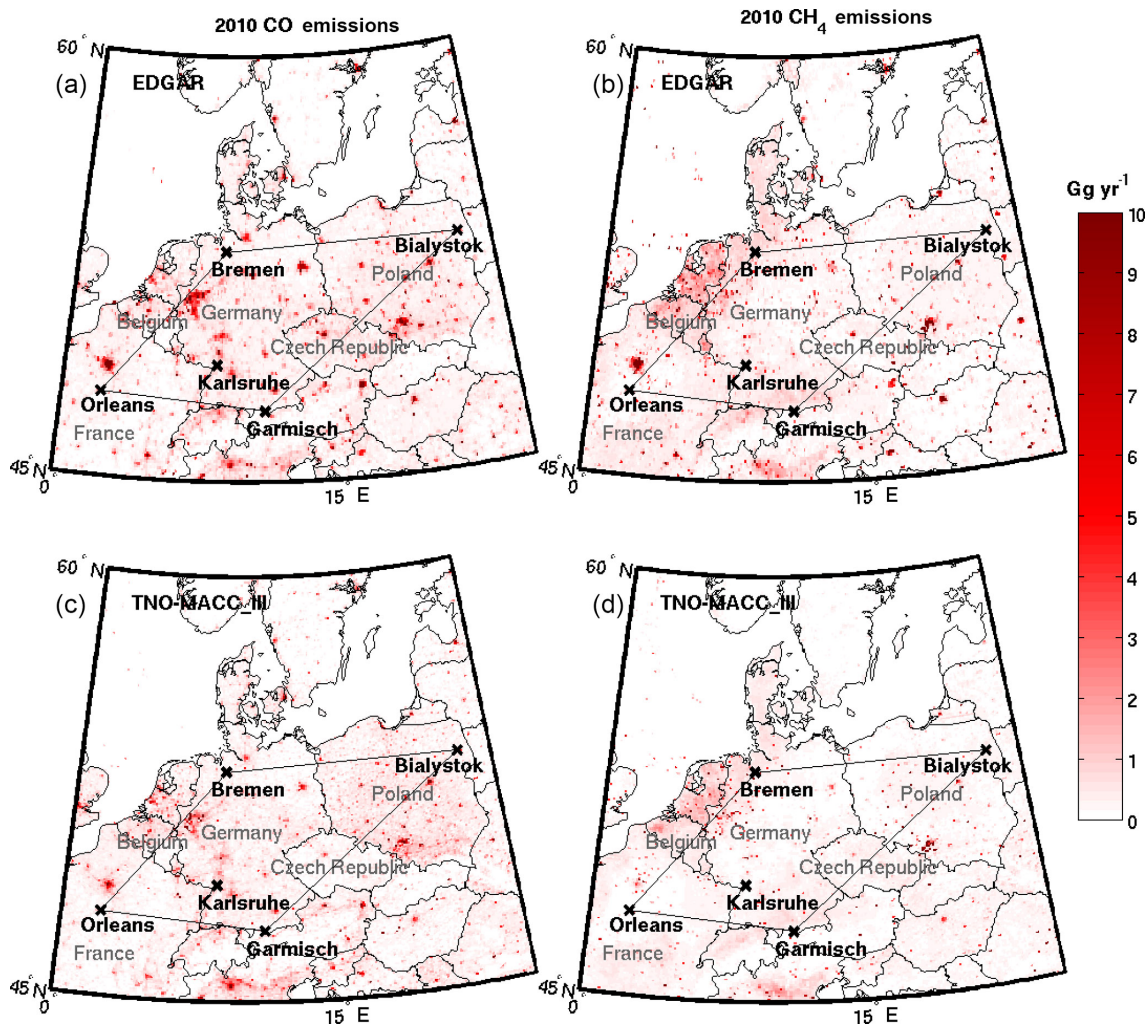


Figure A3. These maps show the inventory emissions for the year 2010 in the study area (delineated by the solid straight lines) and the surrounding region. The TCCON stations are marked with black “x” symbols and labeled in black bold font. The countries intersected by, or contained within, the study area are labeled in grey. The map in (a) shows the EDGAR v4.3.1 emissions inventory for carbon monoxide. The map in (b) shows the EDGAR FT2010 emissions inventory for methane. The map in (c) shows the TNO-MACC_III emissions inventory for carbon monoxide. The map in (d) shows the TNO-MACC_III emissions inventory for methane.

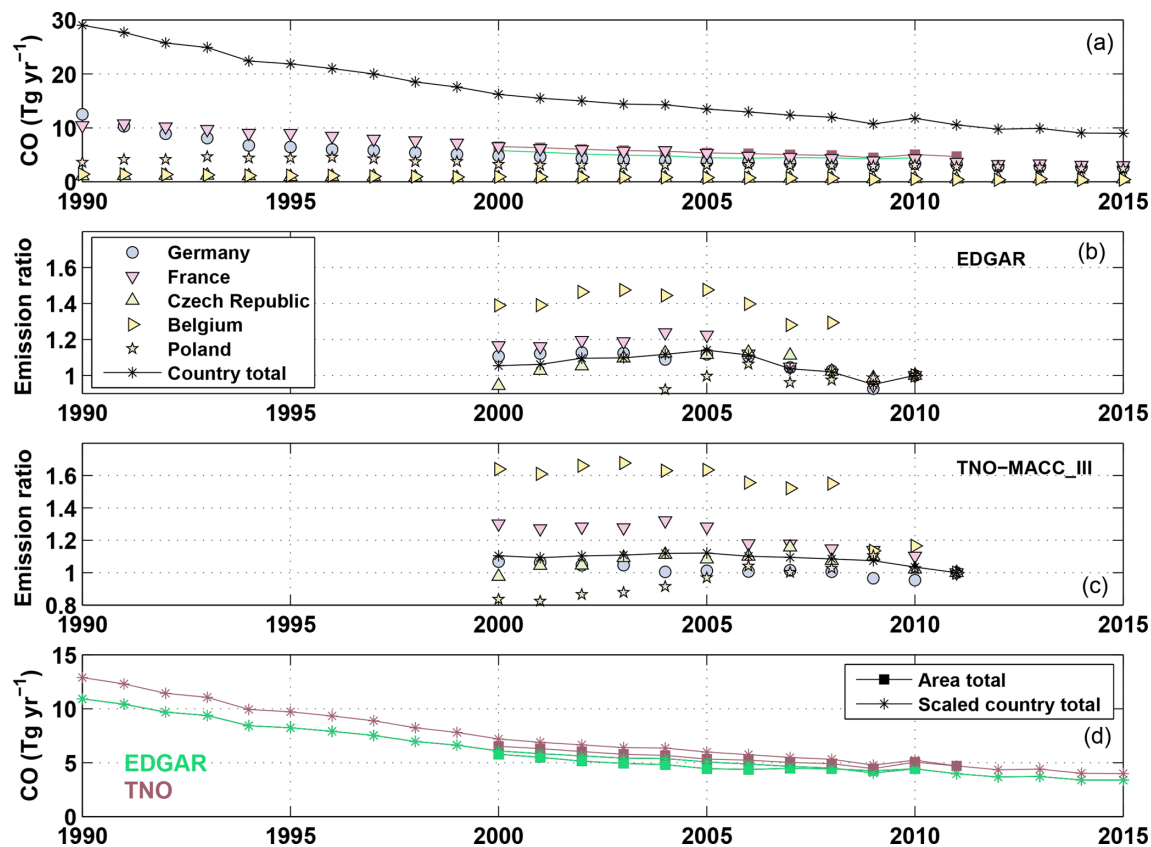


Figure A4. This four-panel plot shows the methodology for scaling the country-level reported emissions of CO to extrapolate the gridded inventory emissions to 2015. Panel (a) shows the CO emissions reported by the European Environment Agency (EEA) for the countries contained within the study area (Germany, France, Czech Republic, Belgium, Luxembourg, and Poland). The black stars with a joining line represent the summed total from the five countries. The EDGAR (green) and TNO-MACC_III (orange) inventories summed within the study area are plotted with squares joined by solid lines. Panel (b) shows the ratio between the individual country totals and the EDGAR area total, normalized to produce an emission ratio of 1 in 2010. The quantity with the least interannual variability in the ratio is from the country total (black stars with line). Panel (c) shows the ratio between the individual country totals and the TNO-MACC_III area total, normalized to produce an emission ratio of 1 in 2011. The quantity with the least interannual variability in the ratio is, again, from the country total. Panel (d) shows the scaled country total, normalized to produce the EDGAR CO emissions for 2010 and the TNO-MACC_III CO emissions for 2011. This permits us to compute a sensible emission for the study area through to 2015.

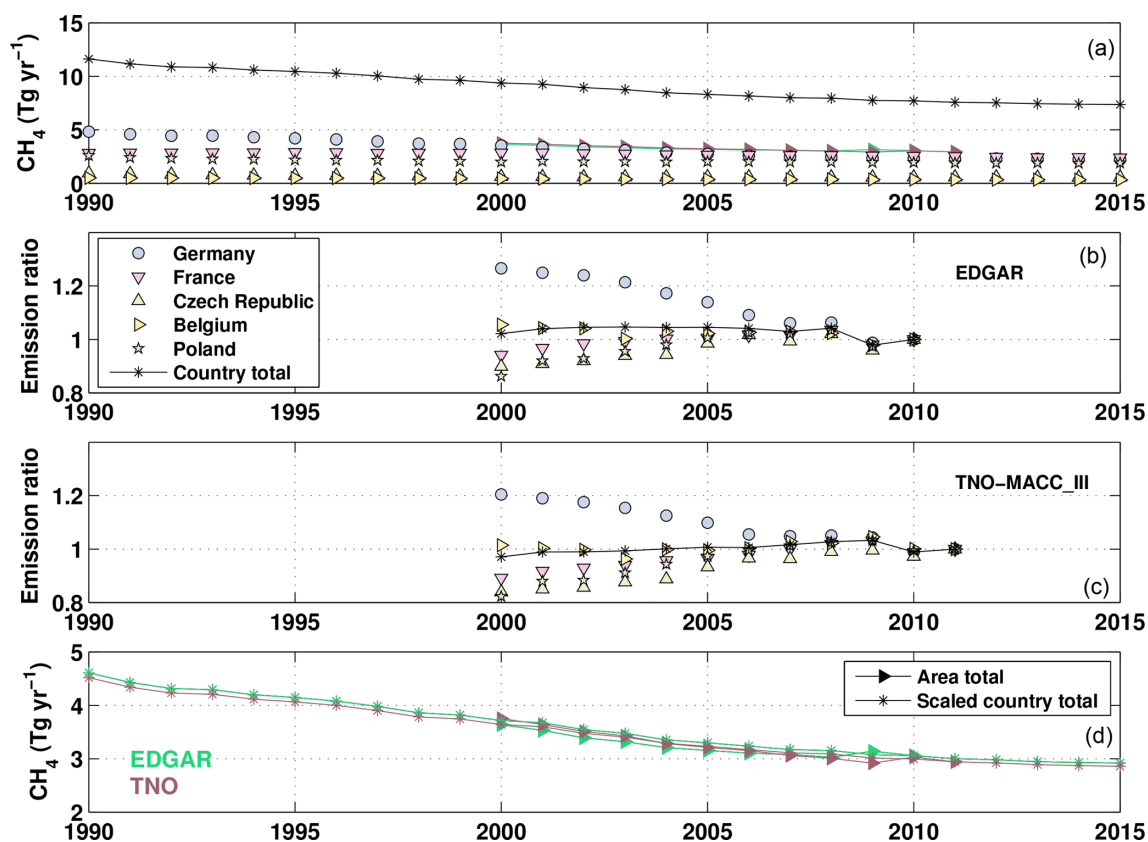


Figure A5. This four-panel plot shows the methodology for scaling the country-level emissions of CH₄ reported to the UNFCCC to extrapolate the gridded inventory emissions to 2015. The panels and symbols follow the same description as in Fig. A4.

Appendix B: Transport time between stations

Figure A1 shows the annual change in monthly mean climatological wind speeds from the NCEP/NCAR reanalysis (Kalnay et al., 1996). These are interpolated to surface pressure and 850 hPa pressures (~ 1500 m geopotential height) from model (sigma) surfaces and cover from January 1948 through March 2017. Vertical mixing into the boundary layer occurs on the timescale of a day or two (Jacob, 1999), and thus the relevant wind speed is between the surface and 850 hPa. The annual mean surface wind speed is 6 km h^{-1} , which gives a mean transit time between Orléans and Białystok of 11 days. The annual mean 850 hPa winds are 17 km h^{-1} , which give a shorter mean transit time between Orléans and Białystok of 4 days.

To test whether the transport time impacts the anomalies, we computed the slopes for time lags between sites of 0–14 days. Figure A2 shows a small change in anomaly slope as a function of the lag used to calculate the anomalies. This figure shows that the transport time between TCCON stations is of negligible importance to the slopes and lends weight to the decision to compute anomalies from data recorded at two TCCON stations on the same day.

Appendix C: Computing study area emissions from the inventories

The study area emissions for 2010 are shown in Fig. A3. We define the study area as the area bounded by the TCCON stations at (clockwise from the west) Orléans, Bremen, Białystok, and Garmisch-Partenkirchen, which is marked by the black lines in the figure. To compute the emissions from the study area, the grid points intersected by and contained within the solid black lines are summed for each year. The EDGAR v4.3.1_v2 emissions inventory for CO and FT2010 inventory for CH₄ provide estimates for years 2000–2010. The TNO-MACC_III inventory provides emissions estimates for both CO and CH₄ for the years 2000–2011.

Appendix D: Projecting inventory emissions beyond 2010

Using data from the European Environment Agency National Database (European Environment Agency, 2016), we extrapolate the inventory CO and CH₄ emissions for the study area through 2015. This is done by summing the total emissions for the five countries that are intersected by the study area (France, Belgium, Germany, Poland, Luxembourg, Czech Republic), and normalizing the emissions to the last year of the inventory (2010 for EDGAR, 2011 for TNO-MACC_III). Figures A4 and A5 show the process for the EDGAR and TNO-MACC_III CO and CH₄ emissions, respectively.

Figure A4a shows the reported country-level emissions for the years 1990–2015, their sum (black stars), and the sum of

the inventory emissions for the years available (2000–2010 for EDGAR; 2000–2011 for TNO-MACC_III) in squares. Figure A4b–c show the ratio of the country-level emissions to the area emissions, normalized to 1 for the last year available in the inventory. These panels show that the ratio of the summed country total emissions to the emissions from the area of interest is less variable from year to year than the emissions reported for individual countries. Thus, we choose to extrapolate the area emissions using the country total emissions, scaled to the last year of the inventory for the study area.

Figure A4d shows the results of using a single scaling factor to estimate the study area emissions from the country-level emissions for each year. We use the summed study area emissions for the years available, and the extrapolated emissions through 2015 for subsequent analysis (e.g., Figs. 3 and 6).

Author contributions. DW designed the study, performed the analysis, and wrote the paper. DBAJ ran the GEOS-Chem model, supported by the CO and CH₄ work of JAF and JDM. JK and HD-*vdG* provided the TNO-MACC_III inventory. GCT helped refine the data analysis methodology. NMD, FH, JN, RS, and TW provided TCCON data. All coauthors read and provided feedback on the contents of the paper and helped interpret the results.

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